# grain size the good, the bad ... and the ugly 

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In this talk, we will follow a fictitious rock through an imagined geological cycle:

- sedimentation
- cementation
- grain growth
- dynamic recrystallization
- fragmentation
- healing

... and observe the associated "grain size"


## "grain size" I - sieving



## "grain size" 2 - strewn samples



## "grain size" 3 - intercept length


closes
average transsected particle width

assumption:
(I) grain volume increase $=74 / 26=1.354$

(2) $d_{\text {cemented }}:$ doriginal $=\sqrt[3]{ } 1.354=1.106$

## "grain size" 4 - numerical simulation



## grain g

Chen LQ, Yang V
Computer-simula quenched system order parameter
https://www.youtube.com/watch?v=p0rY2r0E_2k
"Grain size":

- arithmetic mean of 2D diameter
- mean/mode of curve fit


## "grain size" 5 - thin sections



## "grain size" 6 - particle analyzer


$\theta_{c}^{\text {вобой }}$

## volume(\%) vs. 3D size



Richter, B., 2017,The brittle-to-viscous transition in experimentally deformed quartz gouge. Dissertation, Basel University. https://edoc.unibas.ch/57805/

## "grain size" 7 - multi-scale analysis


fractal dimension for 2 or 3 dimensions

$$
D_{2 d}=D_{3 d}-1
$$

Fractal size distributions should span at least 3 orders of magnitude
grain size ...

## why look at grain size?

grain size data carries information
sediments, sands, silts
$\rightarrow$ environment of deposition
statically recrystallized rocks
$\rightarrow$ time and conditions of grain growth
dynamically recrystallized rocks
$\rightarrow$ level of flow stress
crushed rocks, powders
$\rightarrow$ types of fragmentation processes
... etc.

## the size of a grain - a scalar


thin sections:
size $=$ diameter, d , of area
area of circle: $A=\pi \cdot r^{2}$

$$
\begin{array}{r}
\Rightarrow d=2 \cdot 2 \sqrt{ }(\mathrm{~A} / \pi) \\
\mathrm{d}=2 \mathrm{r} \text { (lower case) }
\end{array}
$$

loose grains, particles: size $=$ diameter, D , of volume
volume of sphere: $V=4 \pi / 3 \cdot R^{3}$

$$
\begin{gathered}
\Rightarrow D=2 \cdot 3 \sqrt{ }(3 \mathrm{~V} /(4 \pi)) \\
D=2 R \text { (upper case) }
\end{gathered}
$$

## the (in)famous 'mean grain size'

| arithmetic mean | $\overline{\mathrm{X}}$ | $=\mathrm{I} / \mathrm{n} \cdot \sum \mathrm{x}_{\mathrm{i}}$ |
| ---: | :--- | :--- |
| geometric mean | G | $=\mathrm{n} \sqrt{ } \Pi \mathrm{x}_{\mathrm{i}}$ |
| harmonic mean | H | $=\mathrm{I} /\left(1 / \mathrm{n} \cdot \sum \mathrm{I} / \mathrm{x}_{\mathrm{i}}\right)$ |
|  |  | $=\mathrm{n} / \sum \mathrm{I} / \mathrm{x}_{\mathrm{i}}$ |
| root-mean-square | RMS | $=\sqrt{ }\left(1 / \mathrm{n} \cdot \sum \mathrm{x}_{\mathrm{i}}{ }^{2}\right)$ |

$$
\begin{aligned}
& \Sigma=\text { sum } \\
& \Pi=\text { product } \\
& \mathrm{i}=\mathrm{I}, \ldots \mathrm{n}
\end{aligned}
$$

$$
\text { RMS }>\bar{X} \geq G \geq H
$$

median
mode
$=x_{(n+1) / 2} \quad$ if $n=$ odd
$=\left(x_{n / 2}+x_{n / 2+1}\right) /$ if $n=$ even
$=$ most frequent value

## why 3D ?!

## ... that's why!



## we 'see' 3D, not 2D, modal grain size




| number-weighted |  |
| :--- | :--- |
| 2D mean | $I .31$ |
| 2D st.dev. | 0.43 |


$\neq$ visual impression

volume-weighted
3D mode 1.70
3D st.dev. 0.1I

= visual impression

## finding the mode by curve fitting



Normal curve fit

$$
\frac{1}{\sqrt{2 \pi \sigma^{2}}} \cdot \exp \left(-\frac{(\mu-x)^{2}}{2 \sigma^{2}}\right)
$$

Lognormal curve fit

$$
\frac{1}{x \sqrt{2 \pi \sigma^{2}}} \cdot \exp \left(-\frac{(\ln (x)-\mu)^{2}}{2 \sigma^{2}}\right)
$$

Polynomial curve fit

$$
m_{0}+m_{1} x+m_{2} x^{2}+m_{3} x^{3}+\ldots
$$

in all cases: fit through center of bin!

## "grain size" | sedimentation sieving

## "grain size" I: beach sand



Great Exhibition Bay, NZ
https://www.geological-digressions.com/analysis-of-sediment-grain-size-distributions/
low tide surf zone

coastal foredunes

"Grain size":
Mean $M \varphi=\left(\varphi_{16}+\varphi_{50}+\varphi_{84}\right) / 3$
Sorting $\sigma \varphi=\left(\varphi_{84}-\varphi_{16}\right) / 4+\left(\varphi_{95}-\varphi_{5}\right) / 6.6$ $\varphi$ equivalent to $\underline{\log (3 D ~ d i a m e t e r) ~}$

## from logarithmic to linear


$\Delta \log (D)=$ constant $\Delta D \neq$ constant

vol / $\Delta \mathrm{D}=$ density
blue, green, ... etc. area on $\log (\mathrm{D})$-plot $=$ blue, green,.. etc. area on D-plot
$\operatorname{vol}(D)=$
density function

$\Delta \mathrm{D}=$ constant
$\Sigma$ area of histogram bars $=$ area under curve $=$ constant

## $\varphi$-values... - double trouble





> weight\% vs. $\varphi$ Gaussian normal fit $M \varphi=2.72 \Rightarrow D_{\text {mean }}=0.152 \mathrm{~mm}$
weight\% vs. $\mathrm{D}(\mathrm{mm})$ w\% binwidth corrected overlay = cubic spline fit (10pts)

## converting the data


weight\% vs. D(mm) d from cubic spline fit overlay $=$ cubic spline fit (39pts)

weight\% vs. $D(\mathrm{~mm})$ Gaussian normal fit
$D_{\text {mean }}=0.212 \mathrm{~mm} \quad(\sigma=0.130 \mathrm{~mm})$


weight\% vs. $D(\mathrm{~mm}) \mathrm{d}$ from cubic spline fit overlay = cubic spline fit (10pts)

weight\% vs. $D(\mathrm{~mm})$ Gaussian normal fit
$D_{\text {mean }}=0.154 \mathrm{~mm} \quad(\sigma=0.020 \mathrm{~mm})$

## $\varphi$-derived versus converted


$M \varphi=2.72$
$D_{\text {mean }}=0.30 \mathrm{~mm}$
weight\%.vs. $\varphi$
$\Longrightarrow D_{\text {mean }}=0.15 \mathrm{~mm}$

weight\%.vs. $\varphi$
$\Longrightarrow D_{\text {mean }}=0.24 \mathrm{~mm}$
weight\%.vs.D(mm)
$D_{\text {mean }}=0.15 \mathrm{~mm}$

## $\sqrt[4]{4}$ 논 $D_{\text {mean }}$ from $\varphi$

 depends on width of distribution$D_{\text {mean }}=$ mode of vol\%(D)

## what have we learned?

Results from sieving are difficult to interpret (... unless you are a sedimentologist ...)

To derive a meaningful mean grain size, $\varphi$-values are best converted to vol\% vs. linear size.

Derived $D_{\text {mean }}$ - values depend on standard deviation.

# "grain size" 2 <br> glacial transport strewn samples 

## "grain size" 2: glacigenic sediments



## volume weighting





## volume\%

> $D_{\text {mean }}$ from vol\%(D) $\neq D_{\text {mean }}$ from $h(D)$

Note: $D_{\text {projected }}=$ diameter of area-equivalent circle of projected area

## what have we learned?

Sands and powders are easily analyzed using a scanner.
In this case, the area-equivalent diameters $\mathrm{d}_{\text {equ }}=\mathrm{D}_{\text {equ }}$ represent the diameters of the volume-equivalent spheres $D_{\text {equ. }}$

No conversion from 2D to 3D is necessary
The conversion from $h(D)$ to vol(D) is trivial: $\operatorname{vol}(\mathrm{D})=h(D) \cdot D^{3}$

# "grain size" 3 cementation intercept method 

## "grain size" 3: cemented sandstone


closest packing $=74 \mathrm{vol} \%$

assumption:
(I) grain volume increase $=74 / 26=1.354$
(2) $d_{\text {cemented }}:$ doriginal $=\sqrt[3]{ } 1.354=1.106$

## intercept method - limitations



$$
\text { grainsize }=\frac{\text { length of test line }(\mathrm{L})}{\text { number of transected grains }(\mathrm{N})}
$$

from grain boundaries:
$L=23100 \mu \mathrm{~m}, \mathrm{~N}=22 \mathrm{I}$
size of (grains+cement) $=\mathrm{L} / \mathrm{N}=105 \mu \mathrm{~m}$

for uncemented grains:
??

| "Grain size": |
| :--- |
| mean of intercept lengths |
| (= $\underline{2 D}$ size, no distribution) |
|  |
| does not work for grains in matrix |

## check against digital image analysis


check (I) diameters (long axis fit ellipse)
grains from intercept grains from ellipse fit
$105 \mu \mathrm{~m}$
$110 \mu \mathrm{~m}$

check (2) ratios
long axes of fit ellipse:
grains
$110 \mu \mathrm{~m}$
(grains+cement) $\quad 130 \mu \mathrm{~m}$
(grains+cement) : grains = I. $22 \neq|.| |$
area\%
grains
72.9 vol\%
(grains+cement) 100.0 vol\%
(grains+cement) : grains $=1.37 \approx 1.35$

## what have we learned?

The intercept method is practical and fast - can be done at the microscope - or on un-segmented micrographs, ... but ...

Only mean the arithmetic 2D mean can be calculated.

Cannot be used for grains in matrix.

# "grain size" 4 grain growth 2D experiment 

## "grain size" 4: Ostwald ripening



## grain growth kinetics

Chen LQ, Yang W (I994)
Computer-simulation of the domain dynamics of a quenched system with a large number of nonconserved order parameters-the grain-growth kinetics.
Phys Rev B 50: 15752--I5756
https://www.youtube.com/watch?v=p0rY2r0E_2k

## "Grain size": <br> - arithmetic mean of 2D diameter <br> - mean/mode of curve fit


diameter


5700 sqpx
area

grain growth
$\rightarrow$ increasing average size
$\rightarrow$ increasing spread
in terms of normal distribution:
$\rightarrow$ increasing mean ( $\mu$ )
$\rightarrow$ increasing standard deviation ( $\sigma$ )
$\Rightarrow$ distribution matters

## 2D simulation




## stats $=$ arithmetic mean

statistics for output file d-01.out.txt (data not saved - need to copy from screen):

|  | mean | st.dev. |
| :--- | ---: | ---: |
| statistics of d | 19.76689 | 8.74732 |
| statistics of D | 19.70738 | 8.28300 |
| statistics of V | 27.25517 | 8.04039 |
| statistics of $\mathrm{D}^{*}$ | 19.70738 | 8.28300 |
| statistics of $\mathrm{V} *$ | 27.25517 | 8.04039 |

## should we convert to 3D...?



## what have we learned?

In a fully cemented / fully crystallized rock grain growth has to be volume conserving.

Ostwald ripening is a valid model for such a process: starting with a normally distributed grain size, both the mean and the standard deviation increase with time.

$$
\begin{aligned}
& \text { "grain size" } 5 \\
& \text { dynamic } \\
& \text { recrystallization } \\
& \text { from 2D to 3D }
\end{aligned}
$$

## "grain size" 5: sheared quartzite



Heilbronner, R. \& Kilian, R. (20I7). The grain size(s) of Black Hills Quartzite deformed in the dislocation creep regime. Solid Earth, 8, 107I-1093, 2017, doi.org/I0.5I94/se-8-I07I-2017.


## $2 D$ versus 3D - mean versus mode










## detect second maximum



## what have we learned?

Converting 2D grain size data to 3D is highly recommended!

Volume weighted 3D histograms should be used they are free from sectioning artefacts!

Modal 3D grain size identifies the physically most relevant grain size(s).

## "grain size" 6

## powder <br> particle analyzer

## "grain size" 6: crushed quartz



Range from $0.2 \mu \mathrm{~m}$ up to $300 \mu \mathrm{~m}$ Mean length is about $65 \mu \mathrm{~m}$ and Mode between 90 and $95 \mu \mathrm{~m}$.

[^0]


Richter, B., 2017,The brittle-to-viscous transition in experimentally deformed quartz gouge. Dissertation, Basel University. https://edoc.unibas.ch/57805/

## to get the mean grain size



## $\neq$ the mean of log histograms !


mean long axis $=$
$D_{\text {mean }}=17.5 \mu \mathrm{~m}!!$

"mean" long axis $=49 \mu \mathrm{~m}$

## evaluated from linear data <br> $\Rightarrow D_{\text {mean }}=17.5 \mu \mathrm{~m}$

> | depends on upper and lower bound, |
| :--- |
| only true for $(-0.5 \leq \log (D) \leq 2.3)$ |
| i.e., for $\quad(3 \mu \mathrm{~m} \leq \mathrm{D} \leq 200 \mu \mathrm{~m})!!!$ |

## evaluated from log data

mean value of $\log (D)=1.69$
$\Rightarrow D_{\text {mean }}=10^{1.69}=49 \mu \mathrm{~m}$
modal value of $\log (\mathrm{D})=1.95$
$\Rightarrow D_{\text {mean }}=101.95=90 \mu \mathrm{~m}$

Range from $0.2 \mu \mathrm{~m}$ up to $300 \mu \mathrm{~m}$
Mean length is about $65 \mu \mathrm{~m}$ and
Mode between 90 and $95 \mu \mathrm{~m}$.
Richter, B., 2017,The brittle-to-viscous transition in experimentally deformed quartz gouge. Dissertation, Basel University. https://edoc.unibas.ch/57805/

## intermezzo: fractal size distributions



F number of fragments created
C number of fragments being fragmented
$f=C / F$ fragmentation fraction
$\mathrm{R}_{\mathrm{i}} \quad$ size (diameter) of fragment
$\mathrm{N}_{\mathrm{i}}$ number of cracked fragments

Fractal dimension

$$
D=\frac{\log \left(N_{i+1} / N_{i}\right)}{\log \left(R_{i} / R_{i+1}\right)}
$$

published example:
The number $N$ of fragments with cube root of volume greater than $r$ is given as a function of $r$ for broken coal(Bennett, 1936), broken granite from a 61 kt underground nuclear detonation (Schoutens, 1979), and impact ejecta due to a $2.6 \mathrm{~km} \mathrm{~s}^{-1}$ polycarbonate projectile impacting on basalt (Fujiwara et al., 1977). The best-fit fractal distribution from (2.6) is shown for each data set.

Example:

$$
\begin{array}{ll}
\mathrm{F}=8 & \mathrm{R}_{1}=\mathrm{R}_{0} / 2 \quad N_{1}=6 \\
C=6 & R_{2}=R_{0} / 4 \quad N_{2}=36 \\
\mathrm{f}=6 / 8
\end{array}
$$

## maximum value for $D=3.00$ - why ?

map views of cube:

$D=0.00$
$D=1.00$
$D=2.58$
$D=3.00$
The fractal distribution $N_{i}\left(R_{i}\right)$ is characterised by a constant ratio $D=\frac{\log \left(N_{i+1} / N_{i}\right)}{\log \left(R_{i} / R_{i+1}\right)} \quad(0 \leq D \leq 3.00)$
$N_{i+1} / N_{1}=$ frequency ratio of smaller to larger grain size $R_{i} / R_{i+1}=$ size ratio of larger to smaller grain size

At the maximum vale of $D=3.0$, the 2 -dimensional fracture surface (grain boundary surface) is completely room-filling, and thus itself a 3-d volume. A higher value than $D=3$ cannot be attained by this process of fragmentation

## more from the fractal world



## $D_{2 \mathrm{~d}}, D_{3 \mathrm{~d}}, E$ from $N / R$ or $\log (N) / \log (R)$

powerlaw fit to linear data
exponent = D
$D_{3 d}$
$N=R^{\left(-D_{38}\right)}$
$D_{2 d}=D_{3 d}-1$
$N=R^{\left(-D_{2 d}\right)}$
$E=3-D_{3 d}$
$N=R^{(E)}$




linear fit to log data
slope $=D$
$D_{3 d}$
$\log N=-D_{3 d} \cdot \log R$
$D_{2 d}=D_{3 d}-1$
$\log N=-D_{2 d} \cdot \log R$
$E=3-D_{3 d}$
$\log N=E \cdot \log R$

## beware: fractal $\neq$ modal distribution


characteristics:

- fractal dimension D
= grain size ratio
- unbounded:
- no minimum, no maximum
- no mean or mode

characteristics:
- moment of central tendency
$=$ most significant grain size
- mean or mode of distribution
- bounded:
- total (area under curve) $=100 \%$


## ... returning to the talk


what was said in the talk:

- you should correct analyzer vol\% to account for increasing bin width with size: vol $_{\text {corr }}=$ vol\% / bin width
- after calculating N\% from volcorr, N\%
keeping this in mind:
 was plotted versus $\mathrm{D}(\mu \mathrm{m})$ and the powerlaw fit yielded $D_{3 d}>3.0$
how this was explained in the talk:
- the processes of hammering and pestling do not correspond to fractal fragmentation


## ...doing it right



Therefore
I. Clean original data: Plot $=$ vol\% vs. $\log (\mathrm{d})$
2. Convert to linear bin size: $\mathrm{d}=10^{\log (\mathrm{d})}$
Plot $=$ vol\% vs. $\mathrm{d}(\mu \mathrm{m})$
;3. Correct bin width: volco $=$ vol / in width $*$ )
3. Directly
convert vol\% $\rightarrow \mathrm{N} \%$
no\% = vol\% / d ${ }^{3}$
Plot $=\mathrm{N} \%$ vs. $\mathrm{d}(\mu \mathrm{m})$
4. Plot $\mathrm{N} \%$ vs. $\mathrm{d}(\mu \mathrm{m})$ on log-log Fit power-law to full data
5. Fit power-law to cropped ( $1 \mu \mathrm{~m} \leq \mathrm{d} \leq 100 \mu \mathrm{~m}$ )

## what does $\mathrm{D}_{3 \mathrm{~d}}$ mean?

The crystal fragments are broken into small pieces with a hammer and screened with a $100-\mu \mathrm{m}$ sieve.
The coarser fraction is repeatedly pestled and sieved until the overall grain size is less than $100 \mu \mathrm{~m}$.
Bettina Richter (2017) PhD thesis, Basel University



how this is explained:

- the processes of hammering and pestling generate grain size distributions with
$\mathrm{D}_{3 \mathrm{~d}}<3.00$
i.e., compatible with fractal fragmentation processes
- $D_{3 d}=2.24$
(fragmentation fraction $\approx 5 / 8$ ) for grains $>1 \mu \mathrm{~m}$
- $\mathrm{D}_{3 \mathrm{~d}}=1.1 \mathrm{I}$
(fragmentation fraction $\approx 2 / 8$ )
for grains $<2 \mu \mathrm{~m}$
(below grinding limit)


## what have we learned?

Data from particle analyzers are particularly prone to misinterpretation.

The mean of a fractal distribution is quite meaningless.
Rather, convert analyzer data to linear histograms of volume density versus linear size...
... and check on a log-log plot of N.vs.size:
if the slope, D , of the powerlaw fit is a straight line, and if ( $0 \leq-\mathrm{D} \leq 3$ ), the distribution may be due to fractal fragmentation.

# "grain size" 7 <br> friction \& healing fractal dimension 

## "grain size" 7: brittle fault rocks






fractal dimension for 2 or 3 dimensions

$$
D_{2 d}=D_{3 d}-1
$$

... but fractal size distributions should span at least 3 orders of magnitude

## comparing grain size distributions




2

$y=$ number (\%)
$\mathrm{x}=$ bins of $3 \mathrm{D} \mathrm{d}(\mathrm{mm})$


all plotted as vol\% vs. linear D


## typical data ranges




Keulen, N., Heilbronner, R., Stünitz, H., Boullier, A.-M. and Ito, H. (2007). Grain size distributions of fault rocks: a comparison between experimentally and naturally deformed granitoids, J. Struct. Geol., 29, I282-I 300, doi: 10.1016/j.jsg.2007.04.003.

## the 'universal' fractal dimension



Keulen, N., Heilbronner, R., Stünitz, H., Boullier, A.-M., Ito, H. (2007). Grain size distributions of fault rocks: a comparison between experimentally and naturally deformed granitoids, Journal of Structural Geology, 29, I282-I 300, doi: $10.1016 / j . j s g .2007 .04 .003$.

intermediate

gouge

$$
\begin{array}{|l|l|}
D_{3 \mathrm{~d}} \text { cracked } & =\sim 2.5 \\
D_{3 \mathrm{~d}} \text { gouge } & =3.0-3.2!! \\
D_{\text {<grinding limit }} & =1.8-2.0
\end{array}
$$



## $D_{3 d}$ gouge $\neq f($ displacement $)$


high velocity friction experiments (rotary shear apparatus)


Stünitz, H., Keulen, N., Hirose,T., Heilbronner, R. (20IO). Grain size distribution and microstructures of experimentally sheared granitoid gouge at coseismic slip rates - criteria to distinguish seismic and aseismic faults? J. Structural Geology, 32, 59-69, doi: $10.1016 / \mathrm{j} . \mathrm{jsg}$.2009.08.002

## experimental and natural fault rocks

## experimentally produced fault rock $\longleftrightarrow$ naturally produced fault rock



Keulen, N., Heilbronner, R., Stünitz, H., Boullier, A.-M., Ito, H. (2007). Grain size distributions of fault rocks: a comparison between experimentally and naturally deformed granitoids, Journal of Structural Geology, 29, I282-1300, doi:10.1016/j.jsg.2007.04.003.


## rupture - faulting - healing

## deformation and healing experiments


fresh



healed

"A hydrostatic healing law for qtz and fs: $\Delta D(t)=D(t)-D_{f}=A \cdot e^{(-\lambda t)}$,
$\Rightarrow$ Healing of monomineralic gouge:
in $\sim 1$ year at $\mathrm{T}=100^{\circ}-200^{\circ} \mathrm{C}$."

## Nojima Fault

natural fault rocks

healed



Keulen, N., Stünitz, H., and Heilbronner, R. (2008). Healing microstructures of experimental and natural fault gouge. Journal of Geophysical Research-Solid Earth II3.

## what have we learned ?

The fractal dimension of freshly fragmented rocks seems to be a 'universal': $\mathrm{D}_{3 \mathrm{~d}}=2.58$.

Mature gouge is 'supra-fractal' with a saturation values of $D_{3 d}>3.00$, indicating the contribution of nonfractal processes (spalling, abrasion), and is independent of the amount of displacement.

Healing of monomineralic fault rocks (gouge)
yields $D_{3 \mathrm{~d}}=2.58$; healing is very fast (on the order of years); for polymineralic rocks $\mathrm{D}_{3 \mathrm{~d}}$ remains $>3$.

## in summary ...

we have considered ...
I. what $\varphi$-values mean in the physical (linear) world
2. the usefulness of a flatbed scanner
3. the fast and easy intercept method
4. if $d_{\text {mean }}$ from 2D simulations can be extrapolated to 3D
5. how to convert $\mathrm{d}_{\text {mean }}$ from 2D sections to $\mathrm{D}_{\text {mean }}$ in 3D
6. how to derive fractal dimensions from particle analyzers
7. that fractal grain size distributions have no mean

## ... and we found that

for any given distribution of grains ...
the arithmetic mean of $h\left(d_{\text {equ }}\right)$
$\neq$ mean of $h\left(D_{\text {equ }}\right)$
$\neq$ mode of vol\%( $\left.D_{\text {equ }}\right)$
$\neq M \varphi$
$\neq$ mean of vol\%( $\left.\log \left(D_{\text {equ }}\right)\right)$
$\neq$... etc.
see data from:
image analysis
scanner
stripstar
sieving
particle analyzer
$\Rightarrow$ ask yourself:
which "grain size" you need to know

## grain size 3D, 2D

## and fractal

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Martin-Luther-Universität Halle, 23. Januar 2023



[^0]:    "Grain size":
    mean or mode of $\log$ (3D size) (e.g., long axes of particles)

    - arithmetic mean
    - mean/mode of curve fits

